

Introducing the 'ABC model' for the efficient study of convective-scale data assimilation techniques

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Convective-Scale NWP (km grid scale)

Some challenges compared to large-scale NWP

- Dynamics:
 - Highly non-linear, 3D, compressible flow with short predictability timescale.
 - Acoustic modes appear in the solution to model equations (clever solvers needed).
 - Near-hydrostatic balance (non-convecting regions).
 - Non-hydrostatic flow (convective regions).
 - Significant ageostrophic component on the small scale.
 - Small-scale quantities become important (e.g. vertical motion, hydrometeors).
- Important observations:
 - Radar reflectivity, IR/MW sounders/imagers.
 - High-density observations (correlated errors).

The 'ABC' Simplified Model

Research brief for simplified model to facilitate convective-scale DA research

- Starting point: the 3D compressible Euler equations.
- Non-hydrostatic, compressible.
- Unbalanced on small scales, balanced on large scales.
- Conserve mass and energy:

$$\frac{d}{dt} \int \rho(x, z, t) dx dz = 0 \quad \frac{d}{dt} \int E(x, z, t) dx dz = 0.$$

- Cheap to run:
 - Integrated with an explicit numerical scheme.
 - 2D (longitude/height domain).
- Dry dynamics (currently).

ABC Model: Tunable Parameters, A , B , C

$$\frac{\partial u}{\partial t} + B\mathbf{u} \cdot \nabla u + C \frac{\partial \tilde{\rho}'}{\partial x} - fv = 0$$

$$\frac{\partial v}{\partial t} + B\mathbf{u} \cdot \nabla v + fu = 0$$

$$\frac{\partial w}{\partial t} + B\mathbf{u} \cdot \nabla w + C \frac{\partial \tilde{\rho}'}{\partial z} - b' = 0$$

$$\frac{\partial \tilde{\rho}'}{\partial t} + B\nabla \cdot (\tilde{\rho}\mathbf{u}) = 0$$

$$\frac{\partial b'}{\partial t} + B\mathbf{u} \cdot \nabla b' + A^2 w = 0$$

$$p' = C\rho_0\tilde{\rho}'$$

- $\mathbf{u} = (u, v, w)$.
- $\tilde{\rho}'$ scaled density pert: $\tilde{\rho}' = (\rho - \rho_0)/\rho_0$.
- b' buoyancy: $b' = (\theta - [\theta_R + \theta_0(z)])g/\theta_R$.
- A pure gravity wave frequency.
- B modulation $0 < B \leq 1$, to reduce acoustic wave frequencies.
- \sqrt{BC} pure acoustic wave speed.

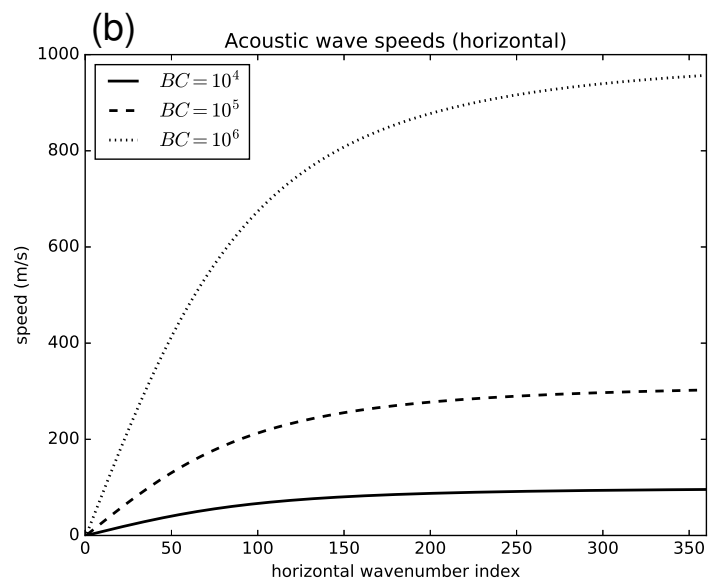
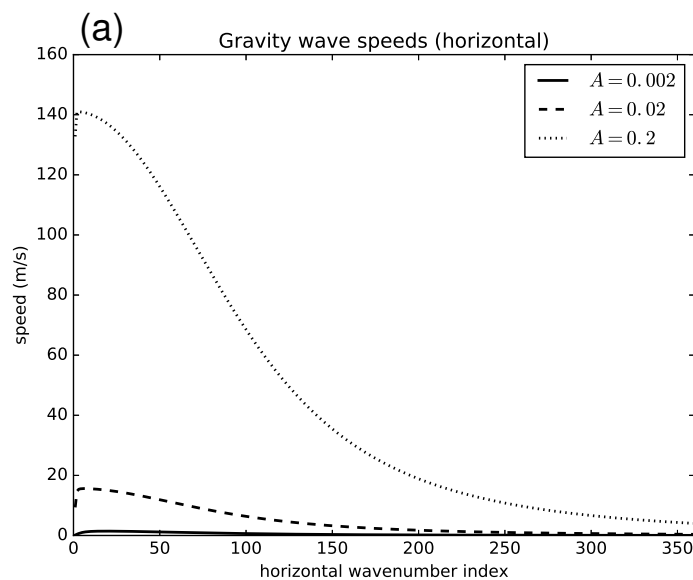
Most experiments: $A = 0.02 \text{ s}^{-1}$, $B = 0.01$, $C = 10000 \text{ m}^2\text{s}^{-2}$.

See Petrie, Bannister, and Cullen (2017).

Model behaviour — linear modes

Horizontal gravity group wave speed

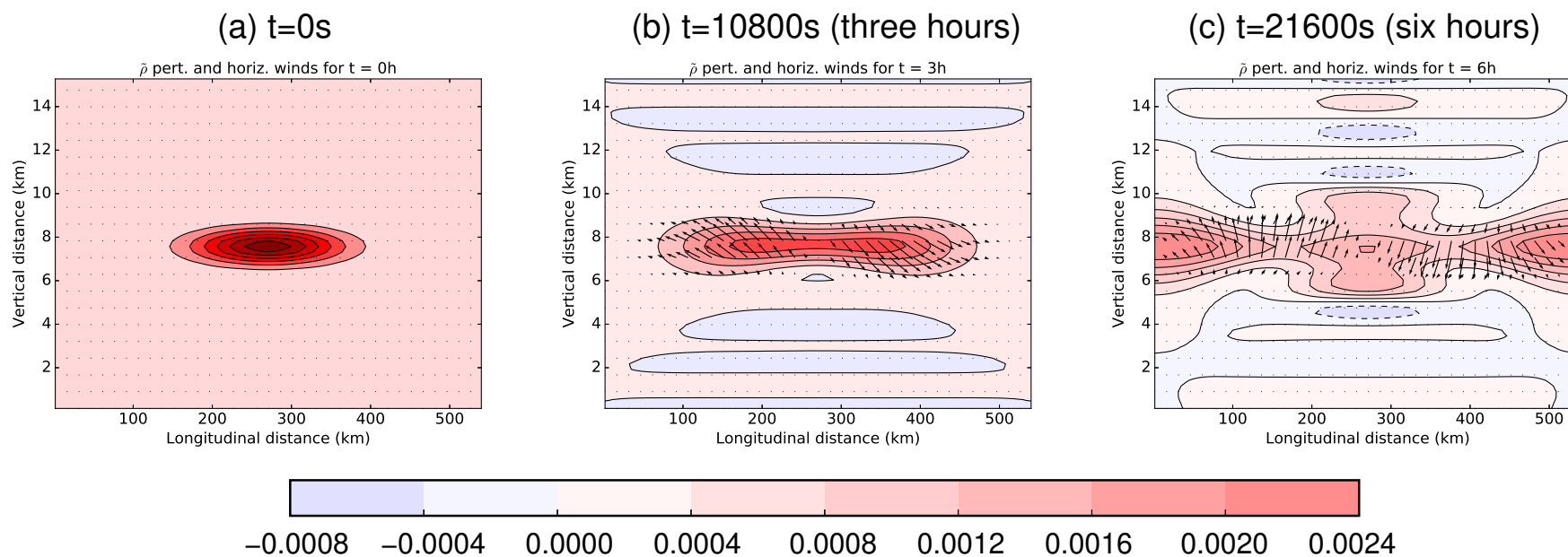
Horizontal acoustic group wave speed



Recall: \sqrt{BC} pure acoustic wave speed

Model behaviour — idealised large-scale

Scaled density pert. ($\tilde{\rho}'$) and horiz. winds (u, v)

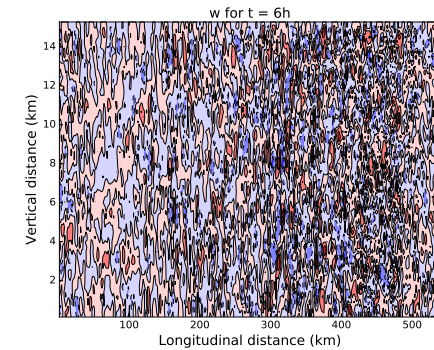
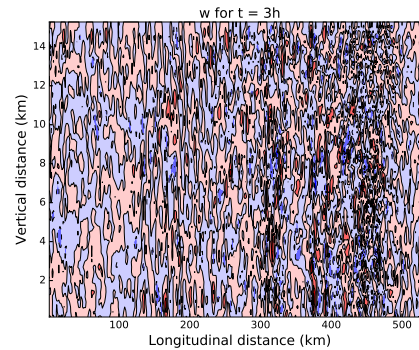
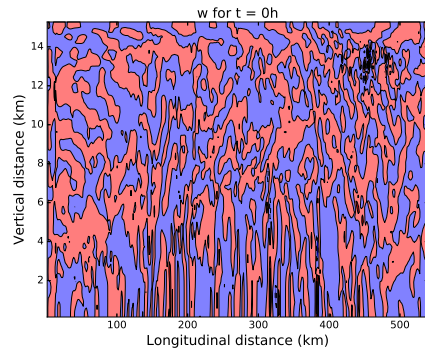


Model behaviour — example integration

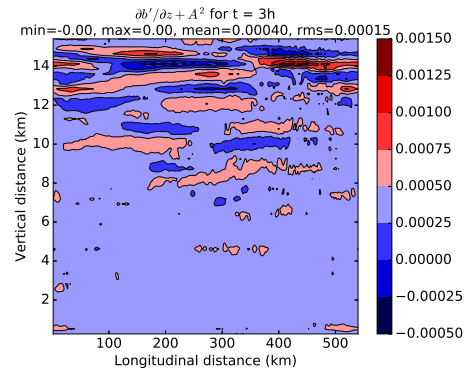
$w(t = 0)$

$w(t = 3 \text{ h})$

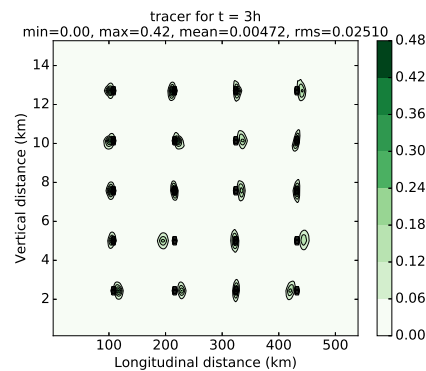
$w(t = 6 \text{ h})$



Effective buoyancy



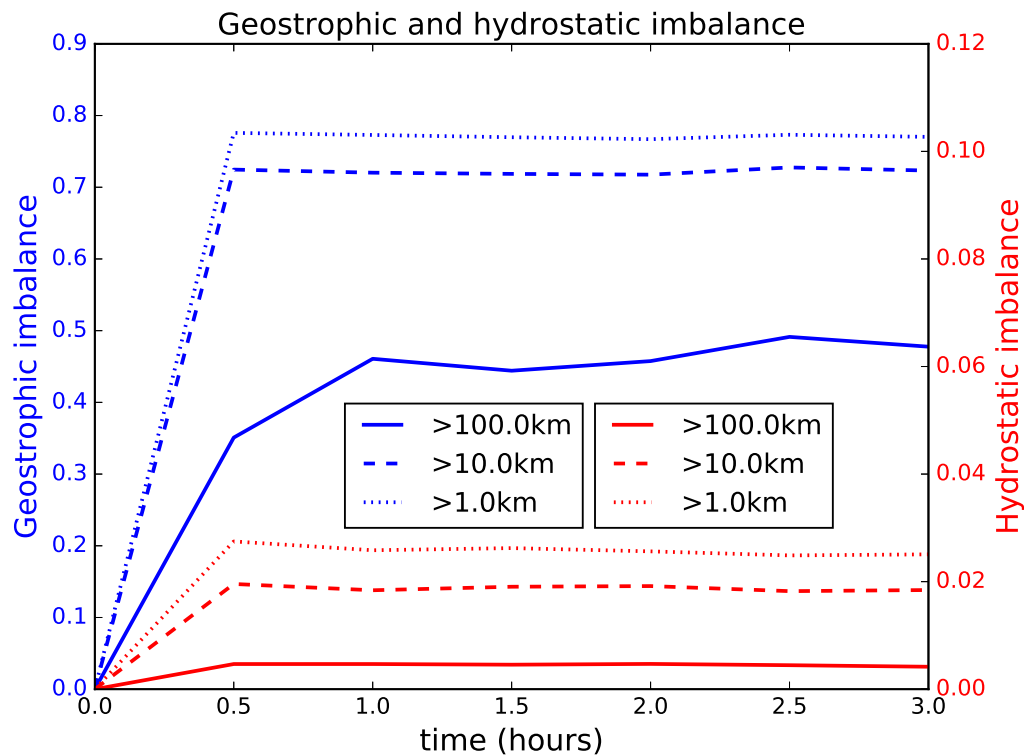
Tracer



$$A^2 + \frac{\partial}{\partial z} [b_0(z) + b'(x, z, t)] =$$

Model behaviour — imbalance

Geostrophic and hydrostatic imbalances



$$\text{geo. imbal.} = \frac{\text{RMS} \left(C \frac{\partial \tilde{\rho}'}{\partial x} - f v \right)}{\text{RMS} \left(C \frac{\partial \tilde{\rho}'}{\partial x} \right) + \text{RMS} \left(f v \right)}$$

$$\text{hydro. imbal.} = \frac{\text{RMS} \left(C \frac{\partial \tilde{\rho}'}{\partial z} - b' \right)}{\text{RMS} \left(C \frac{\partial \tilde{\rho}'}{\partial z} \right) + \text{RMS} \left(b' \right)}$$

Introducing the data assimilation capability

Design of the DA

- Type of DA: variational (currently 3DVar and 3DFGAT implemented)

$$J(\delta\mathbf{x}) = \frac{1}{2}\delta\mathbf{x}^T \mathbf{B}^{-1} \delta\mathbf{x} + \frac{1}{2} (\mathbf{y}_t - \mathcal{H}_t[\mathcal{M}_t(\mathbf{x}^b)] - \mathbf{H}_t \mathbf{M}_t \delta\mathbf{x})^T \mathbf{R}^{-1} (\bullet)$$

$$\delta\mathbf{x} = \mathbf{x} - \mathbf{x}^b.$$

- Background error covariances (\mathbf{B}):
 - Need to be modelled efficiently (control variable transform, $\delta\mathbf{x} = \mathbf{U}\delta\boldsymbol{\chi}$, $\langle \delta\mathbf{x}\delta\mathbf{x}^T \rangle = \mathbf{B}$, $\langle \delta\boldsymbol{\chi}\delta\boldsymbol{\chi}^T \rangle = \mathbf{I}$),
 - Large range of options (e.g. exploiting geophysical balances, options for horizontal and vertical transforms).
- Observations:
 - Flexible observation network (e.g. any quantity, any time/position, allow outlier observations).

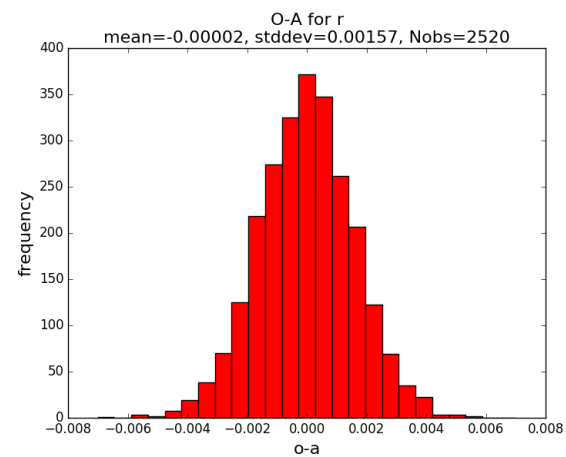
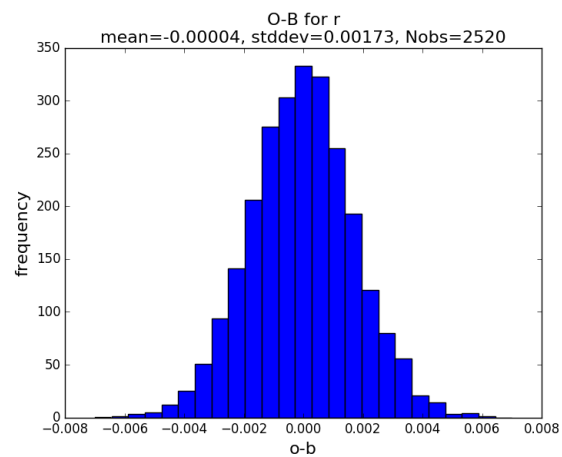
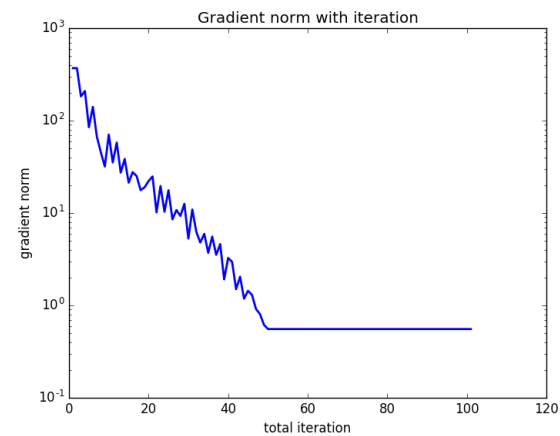
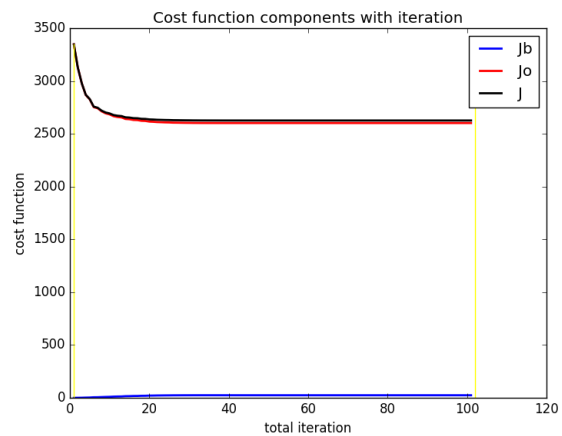
Introducing the data assimilation capability (cont.)

Design of the DA (cont.)

- Software needs to:
 - Calibrate \mathbf{B} -matrix from training data.
 - Test system components (e.g. adjoint tests, inverse tests).
 - Compute implied covariances: $\mathbf{B} = \mathbf{U}\mathbf{U}^T$.
 - Compute raw covariances from training data $\langle \delta\mathbf{x}\delta\mathbf{x}^T \rangle$.
 - Perform 'truth' run, \mathbf{x}^t .
 - Generate \mathbf{x}^b by perturbing \mathbf{x}^t .
 - Generate \mathbf{y}_t by perturbing $\mathcal{H}_t[\mathcal{M}_t(\mathbf{x}^t)]$.
 - Assimilate!
 - Manage cycling.
 - Easy visualisation of assimilation diagnostics.
 - Compare assimilation experiments with different specifications.

Example diagnostics

2520 observations of $\tilde{\rho}'$



Example investigation: geo. bal. in \mathbf{B} -matrix

$$\delta \mathbf{x} = \mathbf{U} \delta \boldsymbol{\chi}$$

$$\begin{pmatrix} \delta \mathbf{u} \\ \delta \mathbf{v} \\ \delta \mathbf{w} \\ \delta \tilde{\boldsymbol{\rho}}' \\ \delta \mathbf{b}' \end{pmatrix} = \begin{pmatrix} \mathbf{0} & \partial_x & \mathbf{0} & \mathbf{0} & \mathbf{0} \\ \partial_y & \mathbf{0} & \mathbf{0} & \mathbf{0} & \mathbf{0} \\ \mathbf{0} & \mathbf{0} & \mathbf{I} & \mathbf{0} & \mathbf{0} \\ \alpha \mathbf{L}^{\text{gb}} & \mathbf{0} & \mathbf{0} & \mathbf{I} & \mathbf{0} \\ \alpha \mathbf{L}^{\text{hb}} \mathbf{L}^{\text{gb}} & \mathbf{0} & \mathbf{0} & \mathbf{L}^{\text{hb}} & \mathbf{I} \end{pmatrix} \begin{pmatrix} \mathbf{U}_s^\psi & \mathbf{0} & \mathbf{0} & \mathbf{0} & \mathbf{0} \\ \mathbf{0} & \mathbf{U}_s^\chi & \mathbf{0} & \mathbf{0} & \mathbf{0} \\ \mathbf{0} & \mathbf{0} & \mathbf{U}_s^w & \mathbf{0} & \mathbf{0} \\ \mathbf{0} & \mathbf{0} & \mathbf{0} & \mathbf{U}_s^{u\tilde{\rho}'} & \mathbf{0} \\ \mathbf{0} & \mathbf{0} & \mathbf{0} & \mathbf{0} & \mathbf{U}_s^{ub'} \end{pmatrix} \begin{pmatrix} \delta \boldsymbol{\chi}^\psi \\ \delta \boldsymbol{\chi}^\chi \\ \delta \boldsymbol{\chi}^w \\ \delta \boldsymbol{\chi}^{u\tilde{\rho}'} \\ \delta \boldsymbol{\chi}^{ub'} \end{pmatrix}$$

$$\text{E.g. } \delta \boldsymbol{\psi} = \mathbf{U}_s^\psi \delta \boldsymbol{\chi}^\psi, \quad \delta \mathbf{u} = \partial_x \delta \boldsymbol{\psi}$$

$$\text{geo. bal. } \delta^b \tilde{\boldsymbol{\rho}}' = \mathbf{L}^{\text{gb}} \delta \boldsymbol{\psi}, \quad \text{hydro. bal. } \delta \mathbf{b}' = \mathbf{L}^{\text{hb}} \delta \tilde{\boldsymbol{\rho}}', \quad \text{geo. bal. switch } \alpha = 0, 1$$

Would like to compute the implied covariances $\mathbf{B} = \mathbf{U}\mathbf{U}^T$.

Example investigation: geo. bal. in \mathbf{B} -matrix (cont.)

$$\mathbf{B} = \mathbf{U}\mathbf{U}^T = \begin{pmatrix} \partial_x \mathbf{B}^x \partial_x^T & \mathbf{0} & \mathbf{0} & \mathbf{0} & \mathbf{0} \\ \mathbf{0} & \partial_y \mathbf{B}^\psi \partial_y^T & \mathbf{0} & \alpha \partial_y \mathbf{B}^\psi \mathbf{L}^{\text{gb}T} & \alpha \partial_y \mathbf{B}^\psi \mathbf{L}^{\text{gb}T} \mathbf{L}^{\text{hb}T} \\ \mathbf{0} & \mathbf{0} & \mathbf{B}^w & \mathbf{0} & \mathbf{0} \\ \mathbf{0} & \alpha \mathbf{L}^{\text{gb}} \mathbf{B}^\psi \partial_y^T & \mathbf{0} & \mathbf{B}^{\tilde{\rho}' } & \mathbf{B}^{\tilde{\rho}' } \mathbf{L}^{\text{hb}T} \\ \mathbf{0} & \alpha \mathbf{L}^{\text{hb}} \mathbf{L}^{\text{gb}} \mathbf{B}^\psi \partial_y^T & \mathbf{0} & \mathbf{L}^{\text{hb}} \mathbf{B}^{\tilde{\rho}' } & \mathbf{L}^{\text{hb}} \mathbf{B}^{\tilde{\rho}' } \mathbf{L}^{\text{hb}T} + \mathbf{B}^{ub'} \end{pmatrix} \begin{matrix} \leftarrow u \\ \leftarrow v \\ \leftarrow w \\ \leftarrow \tilde{\rho}' \\ \leftarrow b' \end{matrix}$$

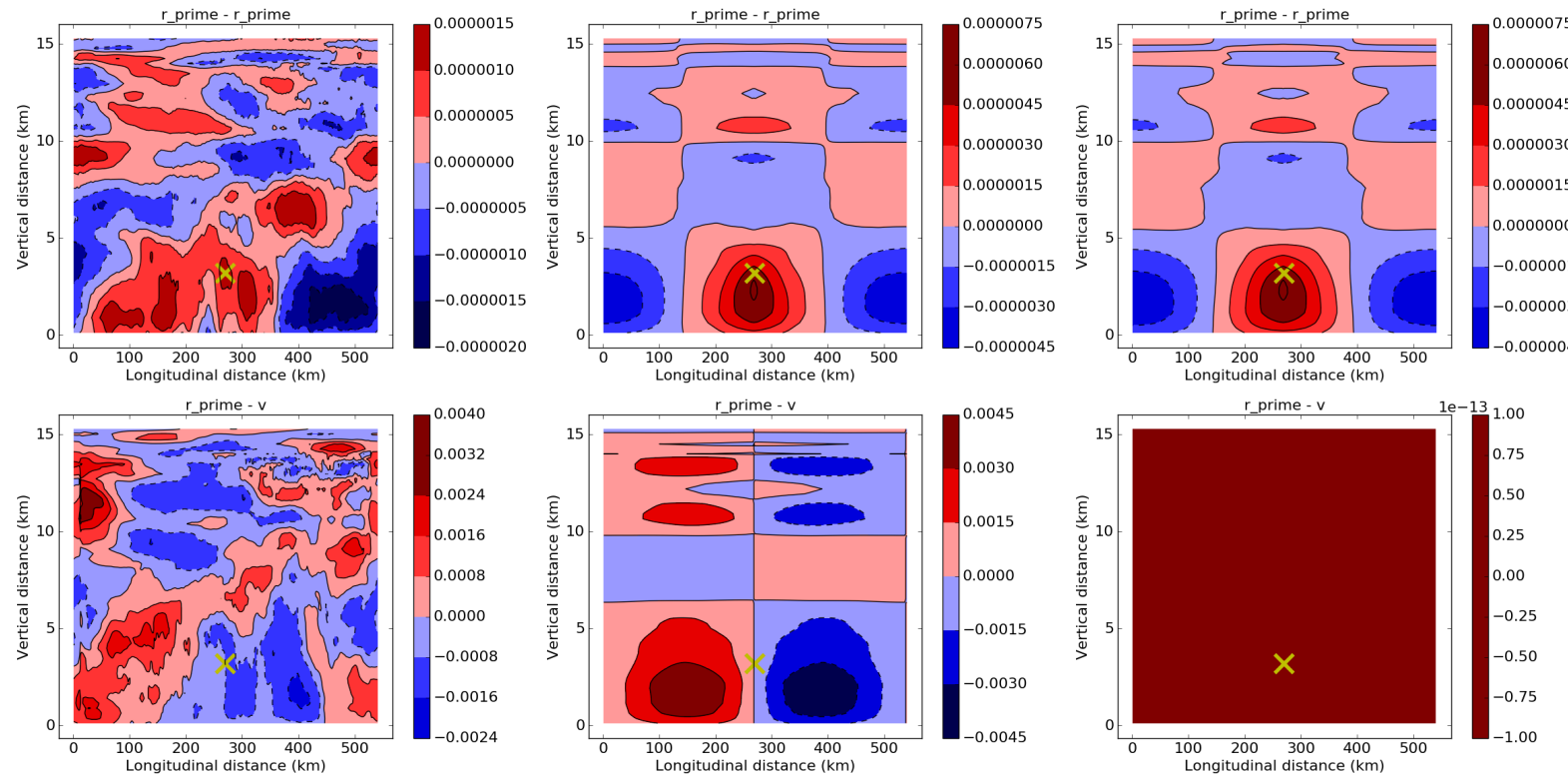
$$\begin{matrix} \uparrow & \uparrow & \uparrow & \uparrow & \uparrow \\ u & v & w & \tilde{\rho}' & b' \end{matrix}$$

where $\mathbf{B}^{\tilde{\rho}' } = \alpha^2 \mathbf{L}^{\text{gb}} \mathbf{B}^\psi \mathbf{L}^{\text{gb}T} + \mathbf{B}^{u\tilde{\rho}' }$, and $\mathbf{B}^\psi = \mathbf{U}_s^\psi \mathbf{U}_s^{\psi T}$, etc.

Example investigation: geo. bal. in B-matrix (cont.)

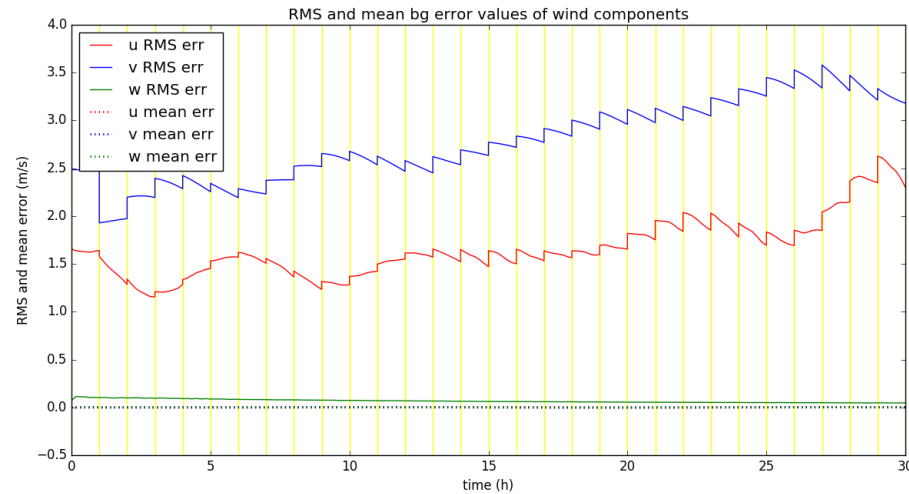
Cov. stats $\langle \delta \mathbf{x} \delta \mathbf{x}^T \rangle$ Implied Cov., $\mathbf{U}\mathbf{U}^T$ (GB+) Implied Cov., $\mathbf{U}\mathbf{U}^T$ (GB-)

$$\tilde{\rho}' - \tilde{\rho}'$$

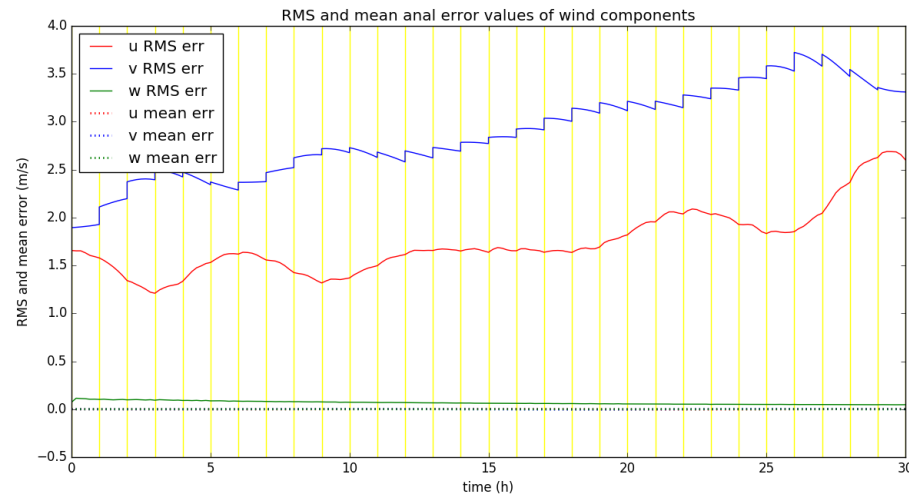


Multi-cycle twin exps (2520 obs of $\tilde{\rho}'$): with geo. bal.

$\alpha = 1$ (GB+), background wind errors

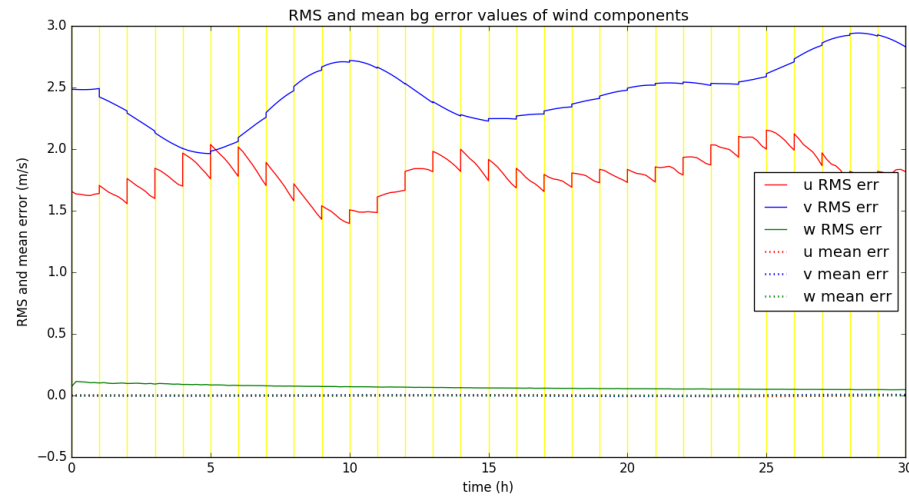


$\alpha = 1$ (GB+), analysis wind errors

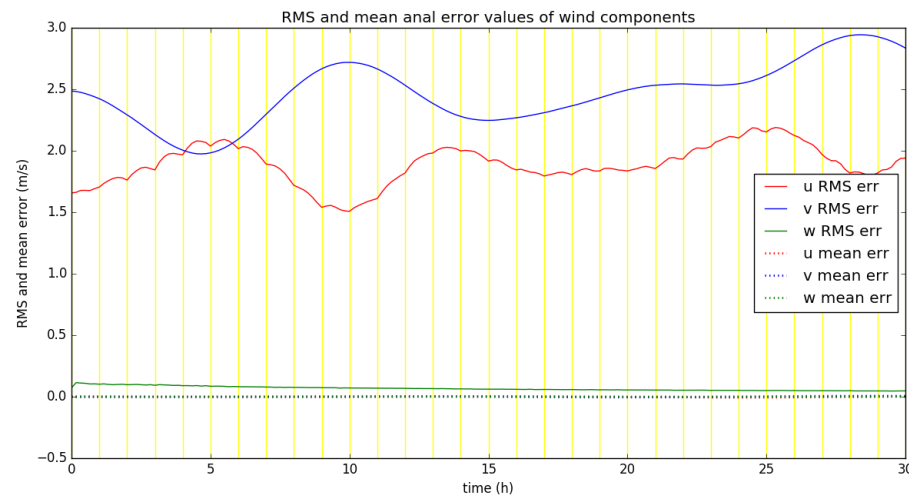


Multi-cycle twin exps (2520 obs of $\tilde{\rho}'$): without geo. bal.

$\alpha = 0$ (GB-), background wind errors



$\alpha = 0$ (GB-), analysis wind errors

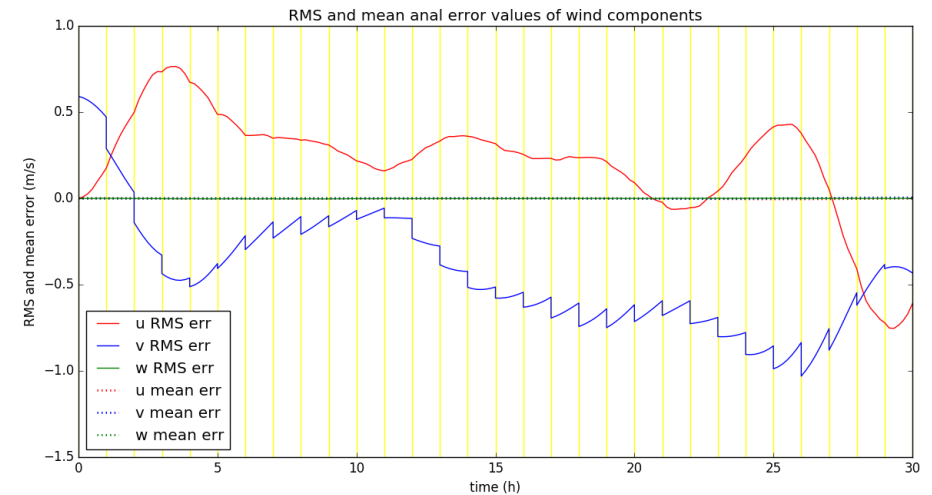


Multi-cycle twin exps (2520 obs of $\tilde{\rho}'$): which is better?

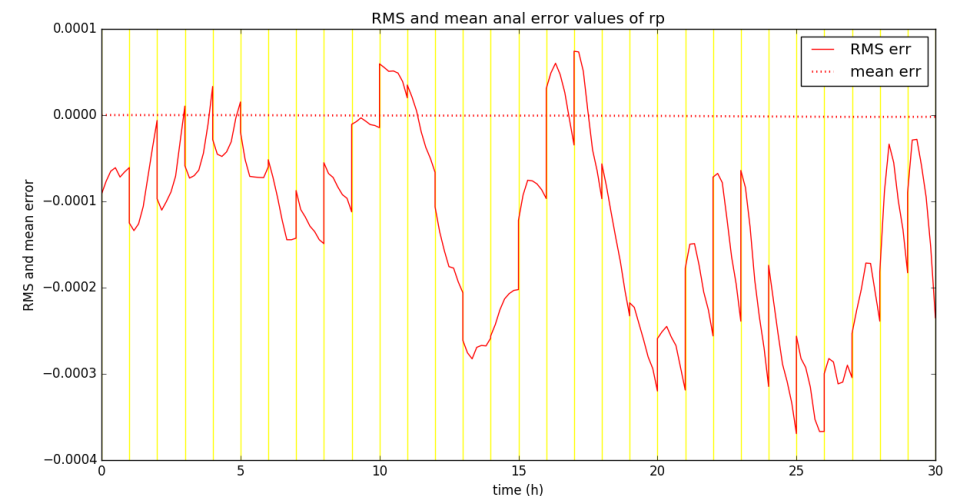
$$\text{diff} = \text{RMSE}(\text{GB-}) - \text{RMSE}(\text{GB+})$$

Negative values indicate that removing geo. bal. operators from \mathbf{B} (i.e. $\alpha = 0$) improves the assimilation.

background wind RMSE differences



background scaled density RMSE differences



Summary

- Have built a **functioning**, **flexible**, and **fast** variational DA system around a simplified 2D convective-scale model.
- **Taster:**
 - Observations of $\tilde{\rho}'$.
 - Switch ON/OFF geostrophic constraint in **B**-matrix.
 - Same initial background state.
 - Cycle length 1 hour, 30 cycles.
 - Twin experiments: know truth.
 - Mixed impact of switching off geostrophic balance – often reduces RMS errors.
- **Future:**
 - Publish system and tests.
 - Investigate other balances.
 - Add: 4D-Var, moisture.
 - Normal mode-based **B**-matrix.
 - Reduced-rank Kalman Filter ...